Numerical Study of Spurious Inertial Modes in Shallow Water Models for a Variable Bathymetry

Gossouhon SITIONON¹, Adama COULIBALY¹, Jérôme Kablan ADOU¹

¹ Université Félix Houphouet Boigny, Abidjan, 22 BP 582 Abidjan 22, UFR Mathématiques et Informatique, Côte d'Ivoire

Correspondence: Gossouhon SITIONON, Université Félix Houphouet Boigny, Abidjan, 22 BP 582 Abidjan 22, UFR Mathématiques et Informatique, Côte d'Ivoire. E-mail: gossouhon.sitionon@gmail.com

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Abstract

In this study we perform a modal analysis of the linear inviscid shallow water equations using a non constant bathymetry, continuous and discontinuous Galerkin approximations. By extracting the discrete eigenvalues of the resulting algebraic linear system written on the form of a generalized eigenvalue / eigenvector problem we first show that the regular variation of the bathymetry does not prevent the presence of spurious inertial modes when centered finite element pairs are used. Secondly, we show that such spurious modes are not present in discontinuous Galerkin discretizations when all variables are approximated in the same descrete space. Such spurious inertial modes have been found very damageable for the quality of inertia-gravity and Rossby modes in ocean modelling.

Keywords: discontinuous Galerkin, mixed finite element, shallow water equations, spurious modes, variable bathymetry

Mathematics Subject Classification: 65M60; 76M10

1. Introduction

For most discretization schemes, the numerical approximation of shallow water models for geophysical flows is a delicate problem. See (Eldred C. & Le Roux D.Y.(2018), Eldred C. & Le Roux D.Y.(2019)). Indeed, the coupling between the continuity and momentum equations usually produces non physical oscillations in the solution, called spurious oscillations. Numerical study of the shallow water equations for finite element pairs that have two velocity components per node brings out three kind of modes namely the physical geostrophic and inertia-gravity modes and non physical spurious inertial modes which appear when the number of discrete momentum equations exceeds, at least by a factor of two, the number of discrete continuity equations. Existence and proportions of theses modes can be found in some papers such as (Le ROUX, 2012) but their authors have only considered the case of constant bathymetry. Indeed, it has been shown that the spurious inertial modes have the frequency $\pm f$, when the Coriolis factor f and the bathymetry H are constant, by using Fourier analysis. Such a Fourier study is not possible when the bathymetry H(x, y) is no longer constant. The first aim of this paper is to examine numerically the existence and proportions among all the discrete frequencies, of the spurious inertial modes for five finite element pairs in the case of a variable bathymetry and compare the results obtained with those of constant bathymetry. The second aim of this paper is to find a method that makes velocity and surface elevation free of spurious inertial modes, namely, the discontinuous Galerkin method.

We first present the model and perform the temporal and space discretizations. The continuous and discontinuous Galerkin formulations are then presented and concluding remarks complete the study.

2. Governing Equations

In this study we focus on the inviscid shallow water (SW) equations linearized around a state of rest, with a variable bathymetry H(x, y) and a constant Coriolis force f. Let Ω be the domain and Σ its boundary. The linearized inviscid shallow water equations in cartesian coordinates are expressed as follows:

$$\widetilde{\mathbf{u}}_t + f\mathbf{k} \times \widetilde{\mathbf{u}} + g \nabla \widetilde{\eta} = 0, \tag{1}$$

$$\widetilde{\eta}_t + \nabla .(H\widetilde{\mathbf{u}}) = 0, \tag{2}$$

where g is the gravitational acceleration, $\widetilde{\mathbf{u}} = (\widetilde{u}, \widetilde{v})$ is the velocity field vector, $\widetilde{\eta}$ is the surface elevation with respect to

the reference level z = 0, **k** is a unit vector in the vertical direction such that $\mathbf{k} \times \widetilde{\mathbf{u}} = (-\widetilde{v}, \widetilde{u})$ and let $\mathbf{x} = (x, y)$. The system (1) –(2) is solved subject to periodic boundary condition to mimic the Fourier analyses of previous studies.

3. Discretization of the Inviscid Linear Shallow-Water Equations

3.1 The Finite Element Method

For a given time step $\Delta t = t_{n+1} - t_n$, $(n = 0; 1; 2; ... and t_0 = 0)$, the time discretization of (1) - (2) is written using a Crank Nicolson scheme. See (Adou, Sitionon & Coulibaly, 2016). By looking for periodic Fourier solutions such as

$$(\widetilde{\mathbf{u}}^{n}(\mathbf{x},t_{n}),\widetilde{\eta}^{n}(\mathbf{x},t_{n})) = (\mathbf{u}(\mathbf{x}),\eta(\mathbf{x}))e^{i\omega t_{n}},$$
(3)

where $\mathbf{u}(\mathbf{x})$ and $\eta(\mathbf{x})$ are the amplitudes of the velocity field and the free surface elevation respectively, with $t_n = n\Delta t$. We let $\alpha = \frac{E-1}{\frac{\Delta t}{2}(E+1)}$ with $E = e^{i\omega\Delta t}$ and we obtain

$$\alpha \mathbf{u} + f \mathbf{k} \times \mathbf{u} + g \nabla \eta = 0, \tag{4}$$

$$\alpha \eta + \nabla .(H\mathbf{u}) = 0. \tag{5}$$

To obtain the weak formulation either the gradient term in (4) or the divergence term in (5) is integrated by parts, which requires two different weak formulations, F1 and F2 respectively, defined as follow.

Formulation F1: Let \mathbf{V}_1 be a subspace of $(H^1(\Omega))^2$, Q_1 be a subspace of $L^2(\Omega)$, $\mathbf{u} \in \mathbf{V}_1$ and $\eta \in Q_1$. The weak formulation is obtained by multiplying (4) by a test function $\phi(\mathbf{x})$ and then multiplying (5) by a test function $\zeta(\mathbf{x})$, with $\phi(\mathbf{x}) \in \mathbf{V}_1$ and $\zeta(\mathbf{x}) \in Q_1$, and integrating over the domain which leads to

$$\alpha \int_{\Omega} \mathbf{u}.\boldsymbol{\phi} d\Omega + f \int_{\Omega} (\mathbf{k} \times \mathbf{u}).\boldsymbol{\phi} d\Omega - g \int_{\Omega} \eta \nabla .\boldsymbol{\phi} d\Omega = 0, \qquad (6)$$

$$\alpha \int_{\Omega} \eta \zeta d\Omega + \int_{\Omega} \nabla (H\mathbf{u}) \zeta d\Omega = 0.$$
⁽⁷⁾

Formulation F2: Let **u** belongs to \mathbf{V}_2 , a subspace of $(L^2(\Omega))^2$ and let η be in a subspace Q_2 of $H^1(\Omega)$. Equation (4) is multiplied by a test function $\varphi(\mathbf{x})$ with $\varphi(\mathbf{x}) \in \mathbf{V}_2$ and equation (5) is multiplyed by a test function $\psi(\mathbf{x})$ belonging to Q_2 and we integrate over the domain Ω to obtain

$$\alpha \int_{\Omega} \mathbf{u}.\boldsymbol{\varphi} d\Omega + f \int_{\Omega} (\mathbf{k} \times \mathbf{u}).\boldsymbol{\varphi} d\Omega + g \int_{\Omega} \boldsymbol{\varphi}.\boldsymbol{\nabla} \eta d\Omega = 0, \qquad (8)$$

$$\alpha \int_{\Omega} \eta \psi d\Omega - \int_{\Omega} H \mathbf{u} \cdot \nabla \psi d\Omega = 0.$$
⁽⁹⁾

The Galerkin method approaches the solutions of (6)-(9) in a finite dimensional subspace. Consider a triangulation T_h , of the polygonal domain Ω , where *h* is the parameter representative of the mesh length which measures the resolution. Let $\mathbf{V}_{j,h}$ and $Q_{j,h}$, be the finite-dimensional subspaces of \mathbf{V}_j and Q_j , respectively, for j = 1, 2. The solution \mathbf{u}_h (resp. η_h) belongs to $\mathbf{V}_{j,h}$ (resp. $\mathbf{Q}_{j,h}$) defined as beeing the set of functions \mathbf{u}_h (resp. η_h) whose restriction on a triangle *K* of T_h belongs to $P_{m_1}(K) \times P_{m_1}(K)$ (resp. $P_{m_2}(K)$), where $P_{m_1}(K)$ (resp. $P_{m_2}(K)$) denotes the set of polynomials of degree m_1 (resp. m_2) defined on *K*.

The introduction of the finite element basis leads to a formulation as in (6)-(9), but with \mathbf{u} , η replaced by the finite element trial functions \mathbf{u}_h , η_h and ϕ , φ , ζ , ψ replaced by the corresponding finite element test functions $\phi_h \in \mathbf{V}_{1,h}$, $\varphi_h \in \mathbf{V}_{2,h}$, $\zeta_h \in Q_{1,h}$ and $\psi_h \in Q_{2,h}$. Further, the solutions \mathbf{u}_h , η_h are expanded over each triangle *K* of T_h in the finite element basis such that

$$\mathbf{u} \simeq \mathbf{u}_h = \sum_{p=1}^{N_s} \mathbf{u}_p \phi_p, \qquad \eta \simeq \eta_h = \sum_{s=1}^{N_v} \eta_s \zeta_s, \tag{10}$$

for the formulation F_1 , and for the formulation F_2 , we obtain

$$\mathbf{u} \simeq \mathbf{u}_h = \sum_{p=1}^{N_s} \mathbf{u}_p \varphi_p, \qquad \eta \simeq \eta_h = \sum_{s=1}^{N_v} \eta_s \psi_s, \tag{11}$$

where η_s and \mathbf{u}_p represent the nodal values of surface elevation and velocity field respectively, while ζ_s , ψ_s and ϕ_p , φ_p are the basic functions of surface elevation and velocity components respectively. N_v is the number of surface elevation nodes and N_s is the number of velocity nodes in the mesh.

Equations (6) - (7) lead to

$$\sum_{K\in T_h} \sum_{p\in I_K} \left[\alpha \mathbf{u}_p + f \mathbf{k} \times \mathbf{u}_p \right] \int_K \phi_p \phi_q d\mathbf{x} - g \sum_{K\in T_h} \sum_{s\in K} \eta_s \int_K \zeta_s \nabla \cdot \phi_q d\mathbf{x} = 0,$$
(12)

$$\sum_{K\in T_h}\sum_{s\in K}\alpha\eta_s\int_K\zeta_s\zeta_r d\mathbf{x} + \sum_{K\in T_h}\sum_{p\in I_K}\mathbf{u}_p \int_K \nabla (H\phi_p)\zeta_r d\mathbf{x} = 0,$$
(13)

and equations (8) -(9) yield

$$\sum_{K\in T_h} \sum_{p\in I_K} \left[\alpha \mathbf{u}_p + f \mathbf{k} \times \mathbf{u}_p \right] \int_K \varphi_p \varphi_q d\mathbf{x} + g \sum_{K\in T_h} \sum_{s\in K} \eta_s \int_K (\nabla \psi_s) \cdot \varphi_q d\mathbf{x} = 0,$$
(14)

$$\sum_{K\in T_h} \sum_{s\in K} \alpha \eta_s \int_K \psi_s \psi_r d\mathbf{x} - \sum_{K\in T_h} \sum_{p\in I_K} \mathbf{u}_p \cdot \int_K H \varphi_p(\nabla \psi_r) d\mathbf{x} = 0,$$
(15)

where $\phi_p = (\phi_p, \phi_p)$, $\varphi_p = (\varphi_p, \varphi_p)$, I_K is the set of velocity nodes belonging to K. Each system in (12) –(15) can be rewritten as a generalized eigenvalues problem

$$\left(\frac{\mathbf{C} \parallel \mathbf{G}}{\mathbf{D} \parallel \mathbf{O}_{1}}\right) \left(\frac{\mathbf{U}}{\boldsymbol{\eta}}\right) = \alpha \left(\frac{-\mathbf{M}^{u} \parallel \mathbf{O}_{3}}{\mathbf{O}_{2} \parallel -\mathbf{M}^{\eta}}\right) \left(\frac{\mathbf{U}}{\boldsymbol{\eta}}\right)$$
(16)

where \mathbf{M}^{u} is the $N_{u} \times N_{u}$ mass matrix for \mathbf{u} , \mathbf{M}^{η} is the $N_{\eta} \times N_{\eta}$ mass matrix for η , \mathbf{D} is the $N_{\eta} \times N_{u}$ divergence matrix, \mathbf{G} is the $N_{u} \times N_{\eta}$ gradient matrix, \mathbf{O}_{1} is the $N_{\eta} \times N_{\eta}$ null matrix, \mathbf{O}_{2} is the $N_{\eta} \times N_{u}$ null matrix, \mathbf{O}_{3} is the $N_{u} \times N_{\eta}$ null matrix, \mathbf{C} is the $N_{u} \times N_{u}$ Coriolis matrix and

$$\mathbf{U} = \left(\begin{array}{ccc} u_1 & v_1 & u_2 & v_2 & \dots & u_{N_u} v_{N_u} \end{array} \right)^T, \quad \boldsymbol{\eta} = \left(\begin{array}{ccc} \eta_1 & \eta_2 & \dots & \eta_{N_\eta} \end{array} \right)^T, \tag{17}$$

with N_u and N_η being the number of velocity and surface elevation nodes, respectively.

3.2 The Discontinuous Galerkin Method

The time discretized equations (4)-(5) are considered and the discontinuous Galerkin method is now employed for the space discretization. Details can be found in (Cokburn, 1997), (Cokburn, 2003) and (Shipton J & Gibson T.H. & Cotter C.J., 2018).

In order to describe the weak formulation, a few notations are first defined. Let $\{T_h\}_{h>0}$ be a partition of the domain Ω into a finite number N_{el} of disjoint open elements K_{el} such that

$$\overline{\Omega} = \bigcup_{i=1}^{N_{el}} \overline{K}_i \text{ and } \overline{K}_{i^+} \cap \overline{K}_{i^-} = \emptyset \text{ for } i^+ \neq i^-,$$
(18)

where *h* is the maximum diameter of the element, $\overline{\Omega}$ is the adhesion of Ω and each element K_i has a Lipschitzian boundary ∂K_i . The mesh length parameter *h* is assumed to be constant. Let Γ be the finite set of *M* inter-element boundaries $\Gamma_{K_i} = \overline{K_{i^+}} \cap \overline{K_{i^-}}$ with $i^+ > i^-$ within the domain, with all possible combinations

$$\overline{\Gamma} = \bigcup_{i=1}^{M} \overline{\Gamma}_{K_i} \text{ and } \overline{\Gamma}_{K_{i^+}} \cap \overline{\Gamma}_{K_{i^+}} = \emptyset \text{ for } i^+ \neq i^-.$$
(19)

Each inter-element boundary $\Gamma_K \in \Gamma$ is associated with a unique fixed unit normal vector denoted by $\mathbf{n}_K = (n_K^x, n_K^y)$. Further, for any function $\zeta \in V_K := H^1(K)$, where ζ represents u, v or η , and for each element K, the trace of ζ on interior edges Γ_K is denoted by ζ^{\pm} , or ζ_L (L for left) and ζ_R (R for right). For $\mathbf{x} \in \Gamma_K$ we have

$$\zeta_g(\mathbf{x}) = \zeta^-(\mathbf{x}) = \lim_{\epsilon \to 0^-} \zeta(\mathbf{x} + \epsilon \mathbf{n}_K) \quad \zeta_d(\mathbf{x}) = \zeta^+(\mathbf{x}) = \lim_{\epsilon \to 0^+} \zeta(\mathbf{x} + \epsilon \mathbf{n}_K). \tag{20}$$

To obtain the weak formulation both gradient term in (4) and divergence term in (5) are integrated by parts, which lead to the weak formulations F3 defined as follow.

Formulation F3: Let V_3 be a subspace of $H^1(K)$, $\forall K \in T_h$, $\mathbf{u} \in V_3 \times V_3$ and $\eta \in V_3$. The weak formulation is obtained by multiplying (4) by a test function $\phi(\mathbf{x}) = (\phi(\mathbf{x}), \phi(\mathbf{x}))$, then multiplying (5) by a test function $\phi(\mathbf{x})$ with $\phi(\mathbf{x}) \in V_3 \times V_3$ and $\phi(\mathbf{x}) \in V_3$, and integrating over the domain leads to

$$\alpha \int_{K} \mathbf{u}.\boldsymbol{\phi} d\mathbf{x} + f \int_{K} (\mathbf{k} \times \mathbf{u}).\boldsymbol{\phi} d\mathbf{x} - g \int_{K} \eta \nabla .\boldsymbol{\phi} d\mathbf{x} + \int_{\partial K} g \eta \boldsymbol{\phi}.\mathbf{n}_{k} ds = 0, \qquad (21)$$

$$\alpha \int_{K} \eta \phi d\mathbf{x} - \int_{K} H \mathbf{u} \cdot \nabla \phi d\mathbf{x} + \int_{\partial K} H \phi \mathbf{u} \cdot \mathbf{n}_{k} ds = 0.$$
(22)

The Discontinuous Galerkin method approaches the solutions of (21)-(22) in a finite dimensional subspace. Let $V_{3,h}$ be the finite-dimensional subspaces of V_3 . The solution \mathbf{u}_h (resp. η_h) belongs to $V_{3,h} \times V_{3,h}$ (resp. $V_{3,h}$) defined as beeing the set of functions \mathbf{u}_h (resp. η_h) whose restriction on a triangle *K* of T_h belongs to $P_{m_1}(K) \times P_{m_1}(K)$ (resp. $P_{m_1}(K)$), where $P_{m_1}(K)$ denotes the set of polynomials of degree m_1 defined on *K*.

The introduction of the finite element basis leads to a formulation as in (21)-(22), but with \mathbf{u} , η replaced by the finite element trial functions \mathbf{u}_h , η_h and $\boldsymbol{\phi}$ replaced by the corresponding finite element test function $\boldsymbol{\phi}_h \in V_{3,h} \times V_{3,h}$, $\phi_h \in V_{3,h}$.

Equations
$$(21) - (22)$$
 lead to

$$\alpha \sum_{K \in T_h} \int_{K} \mathbf{u}_h \cdot \boldsymbol{\phi}_h d\mathbf{x} + f \sum_{K \in T_h} \int_{K} (\mathbf{k} \times \mathbf{u}_h) \cdot \boldsymbol{\phi}_h d\mathbf{x} - g \sum_{K \in T_h} \int_{K} \eta_h \nabla \cdot \boldsymbol{\phi}_h d\mathbf{x} + \sum_{K \in T_h} \int_{\partial K} g \eta^* \boldsymbol{\phi} \cdot \mathbf{n}_K ds = 0,$$
(23)

$$\alpha \sum_{K \in T_h} \int_K \eta_h \phi_h d\mathbf{x} - \sum_{K \in T_h} \int_K H \mathbf{u}_h \cdot \nabla \phi_h d\mathbf{x} + \sum_{K \in T_h} \int_{\partial K} H \phi \mathbf{u}^* \cdot \mathbf{n}_K ds = 0,$$
(24)

where $\phi_h = (\phi_h, \phi_h)$, (\mathbf{u}^*, η^*) denote the numerical trace of (\mathbf{u}, η) on the boundary element ∂K with $\mathbf{u}^* = (u^*, v^*)$. Note that the finite volume scheme is obtained from (23) - (24) when m = 0, and hence the third term in the left hand side of (23) and the second term in the left hand side of (24) are all zeros as the test function belongs to $P_0(K)$. To complete the definition of the approximate solution (u_h, v_h, η_h) , it only remains to choose a unique numerical flux (u^*, v^*, η^*) at the cell interface as to render the method consistent and stable. By using the Rusanov flux which guarentees the stability, we obtain

$$\begin{pmatrix} g\eta^*\mathbf{n}_K\\ H\mathbf{u}^*.\mathbf{n}_K \end{pmatrix} = \begin{pmatrix} g\mathbf{n}_K \{\eta\} - \frac{\sqrt{gH}}{2} [\mathbf{u}]\\ H\mathbf{n}_K.\{\mathbf{u}\} - \frac{\sqrt{gH}}{2} [\eta] \end{pmatrix},$$
(25)

with $\zeta \in \{u, v, \eta\}, [\zeta] = \zeta_R - \zeta_L$ is the jump of ζ and $\{\zeta\} = \frac{\zeta_R + \zeta_L}{2}$ is the mean of ζ . After expanding **u** and η on each triangle *K* of *T_h*, in terms of their respective nodal values and basis functions, we obtain the generalized eigenvalue problem

$$\left(\frac{\mathbf{C} + \mathbf{S}^{u} \| \mathbf{G}}{\mathbf{D} \| \mathbf{S}^{\eta}}\right) \left(\frac{\mathbf{U}}{\eta}\right) = \alpha \left(\frac{-\mathbf{M}^{u} \| \mathbf{O}}{\mathbf{O} \| -\mathbf{M}^{\eta}}\right) \left(\frac{\mathbf{U}}{\eta}\right),$$
(26)

where \mathbf{M}^{u} is the $2N \times 2N$ mass matrix for \mathbf{u} , \mathbf{M}^{η} is the $N \times N$ mass matrix for η , \mathbf{D} is the $N \times 2N$ divergence matrix, \mathbf{G} is the $2N \times N$ gradient matrix, \mathbf{O} is the null matrix, \mathbf{C} is the $2N \times 2N$ Coriolis matrix, \mathbf{S}^{u} is the $2N \times 2N$ stabilization matrix for \mathbf{u} , \mathbf{S}^{η} is the $N \times N$ stabilization matrix for η and

Note that \mathbf{S}^{η} and \mathbf{S}^{η} are the stabilization matrices regrouping the jumps of **u** and η , respectively.

4. Numerical Results

All numerical results are obtained using FreeFem++ of (Hecht, F, 2012). Equations (12) –(15) are solved in a rectangular domain of length L_x and width L_y on a structured mesh, the Coriolis force f being constant. The boundary of the domain is subject to periodic conditions.

This test is performed using finite element pairs $P_0 - P_1$, $P_1^{DG} - P_1$, $P_1^{DG} - P_2$, $P_1^{NC} - P_1$, $P_2 - P_1$, and $P_1^{DG} - P_1^{DG}$. The structured mesh is composed of right isosceles triangles, and *a* and *b* denote the number of segments on the length and on the width of the domain, respectively. Several variable bathymetries have been tested and all of them lead to the same conclusions. We vary the values of *a* and *b* and in each case we compare the number of eigenvalues of (16) and (26) with

the product *ab*. The results presented in this study have been obtained with three variable bathymetries H_1 , H_2 and H_3 defined by (28), (29) and (30) respectively, and shown in Figure 1. The bathymetry H_1 is defined by

$$H_1(x, y) = H_0(1 - d.e^{-c((x-x_0)^2 + (y-y_0)^2)}),$$
(28)

where H_0 is the depth around which the pertubation takes place, and *c* and *d* well chosen constants, with c = 0.02 and d = 0.003, and (x_0, y_0) is the origin of the domain Ω . The bathymetry H_2 on Figure 1 is defined by

$$H_2(x, y) = H_0(1 + d.e^{-c((x-x_0)^2 + (y-y_0)^2)}),$$
(29)

while the bathymetry H_3 reads

$$H_{3}(x,y) = \begin{cases} 1.63, \quad \forall (x,y) \in [-2.39, -1.195] \times [-1.43, 1.43] \\ 0.3x + 1.99, \quad \forall (x,y) \in [-1.195, 1.195] \times [-1.43, 1.43] \\ 2.3485, \quad \forall (x,y) \in [1.195, 2.39] \times [-1.43, 1.43]. \end{cases}$$
(30)



Figure 1. The bathymetries: (left) H_1 , (center) H_2 and (right) H_3

At initial time, an anti-cyclonic velocity field is in geostrophic equilibrium with a prescribed Gaussian distribution of the surface elevation, centered at the origin of the domain. This leads to the following initial conditions

$$\widetilde{\eta}(x, y, t = 0) = \kappa e^{-\sigma((x - x_0)^2 + (y - y_0)^2)},$$
(31)

$$\widetilde{\mathbf{u}}(x, y, t = 0) = \frac{g}{f} \mathbf{k} \times \nabla \widetilde{\eta}(x, y, t = 0),$$
(32)

with κ and σ are well chosen constants, for example $\kappa = 0.02$ and $\sigma = 0.003$. In the model, the equations were adimentionalized with the inertial period T = 104504s, the characteristic time. When computing the eigenvalues of (16), we obtained three kind of values: some values equal to 0, some values equal to ± 7.6178 and the others are such that their absolute value is greater than 7.6178. To get the dimentionalized value, we divide each pulsation value by T. Indeed we have $E = e^{i\omega\Delta t}$, $e^{i\omega\Delta t} = e^{i\omega'\Delta t'}$ and $e^{i\omega\Delta t} = e^{i\omega'\Delta t'}$. So the numerical value ω' computed is such that $\omega' = \omega T$. Thus $7.6178/T \simeq 7.2894 \times 10^{-5}$ and this correspond to $f = 2\Omega_e \sin(\Phi) = \Omega_e = 7.2921 \times 10^{-5} rad.s^{-1}$ for $\Phi = 30^\circ$. The parameter Φ is equal to 30° in the simulation. This implies the following remark.

Remark 1. In the numerical results the values $f = \pm 7.6178$ correspond to the spurious inertial modes, the values 0 correspond to the geostrophic modes and the others values such that their absolute value is greater than 7.6178 correspond to the inertia-gravity modes. We set $f_{sp} = 7.6178$.

We vary the values of *a* and *b* in the range of integers [2; 12]. In each case we compare the number of geostrophic modes $\omega = 0$, inertial modes $\omega = \pm f$ and inertia-gravity modes corresponding to values of ω such that $|\omega| > f_{sp}$ with the product *ab*.

In the continuous case we have one geostrophic mode $\omega = 0$, zero spurious inertial mode $\omega = \pm f$ and 2 inertia-gravity modes $\omega = \pm \sqrt{f^2 + gH(k^2 + l^2)}$, where k and l are the wave numbers in the x- and y-directions, respectively. We denote by GM the geostrophic modes, IGM the inertia-gravity modes, SIM the spurious inertial modes, FE the finite element and RSIM the ratio of spurious inertial modes, namely $RSIM = \frac{N_{SIM}}{N_{GM} + N_{SIM} + N_{IGM}}$, where N_{SIM} , N_{GM} and N_{IGM} are the number of spurious inertial modes, the number of geostrophic modes and the number of inertia-gravity modes respectively.

In order to determine the influence of bathymetry on the existence and proportions of the spurious inertial modes, we firstly compute the number of modes when the bathymetry is constant. We vary the value of (a, b) in any $(a, b) \in [\![2; 12]\!]^2$ and results are shown in Table 1.

Table 1. Results of eigenvalues computed with a constant bathymetry for the $P_2 - P_1$, $P_1^{NC} - P_1$, $P_0 - P_1$, $P_1^{DG} - P_1^{DG}$

Mode	$P_2 - P_1$		$P_{1}^{NC} - P_{1}$		$P_0 - P_1$		$P_{1}^{DG} - P_{1}$		$P_1^{DG} - P_2$		$P_{1}^{DG} - P_{1}^{DG}$	
type	ω	number	ω	number	ω	number	ω	number	ω	number	ω	number
GM	0	ab	0	ab	0	ab	0	ab	0	4ab	0	6ab
SIM	$\pm f_{sp}$	6ab	$\pm f_{sp}$	4ab	$\pm f_{sp}$	2ab	$\pm f_{sp}$	10ab	$\pm f_{sp}$	4ab	$\pm f_{sp}$	0
IGM	ω^{g}	2ab	ω^{g}	2ab	ω^{g}	2ab	ω^{g}	2ab	ω^{g}	8ab	ω^{g}	12ab
RSIM		6/9		4/7		2/5		10/13		4/16		0

Note: ω^g any pulsation ω such that $|\omega| > f_{sp}$.

For all variable bathymetries considered in this study we obtained the same values and numbers of spurious inertial and geostrophic modes. However, if the number of inertia-gravity modes is the same, their value changes only slightly, from one bathymetry to another. For example, when taking (a, b) = (5; 3) we obtain as maximum mode values 8.8942, 8.8943 and 8.8944 for bathymetries H_1 , H_2 and H_3 respectively.

Table 2. The geostrophic, spurious inertial and inertia-gravities modes proportions for the $P_2 - P_1$, $P_1^{NC} - P_1$, $P_0 - P_1$, $P_1^{DG} - P_1$, $P_1^{DG} - P_2$, $P_1^{DG} - P_1^{DG}$ finite element pairs, for any $(a, b) \in [\![2; 12]\!]^2$ with variable bathymetries

Modes	Continuous	$P_2 - P_1$	$P_{1}^{NC} - P_{1}$	$P_0 - P_1$	$P_{1}^{DG} - P_{1}$	$P_{1}^{DG} - P_{2}$	$P_1^{DG} - P_1^{DG}$
GM	0	ab	ab	ab	ab	4ab	6ab
SIM	$\pm f_{sp}$	6ab	4ab	2ab	10 <i>ab</i>	4ab	0
IGM	$ \omega > f_{sp}$	2ab	2ab	2ab	2ab	8ab	12 <i>ab</i>
RSIM		6/9	4/7	2/5	10/13	4/16	0

Further, when using (a, b) = (5; 3) we obtain 90, 60, 30, 150, 60, 0 values equal to $\pm f_{sp}$, namely $f = \pm 7.6178$, corresponding to the number of spurious inertial modes for the finite element pairs $P_2 - P_1$, $P_1^{NC} - P_1$, $P_0 - P_1$, $P_1^{DG} - P_1$, $P_1^{DG} - P_2$, $P_1^{DG} - P_2$, $P_1^{DG} - P_1^{DG}$, respectively.

As a conclusion, the mode proportion results obtained in this section are consistent with those of Table 5 of (Le ROUX, 2012) for all finite element pairs that have two velocity components per node considered in this study. The physical parameters that are employed to discretize the SW equations do not affect the results in these tables for any $(a, b) \in [\![2; 12]\!]^2$. Thus, the results obtained in Table 5 of (Le ROUX, 2012) in the case of a constant bathymetry are the same as for the variable bathymetries used in this study. A regular variation of the bathymetry does not change the proportions of the modes.

We compute the eigenvalues of (26) using the Rusanov flux and we noted that a finite element space discretization using the discontinuous Galerkin method does not reveal any spurious inertial mode for constant and variable bathymetries. This is explained by the fact that the approximation spaces of the velocity field and the elevation of the free surface elevation are the same. Thus the principal responsable of spurious modes seems to be the interaction of different approximation spaces of the finite element pairs that are used.

5. Conclusion

The study of the generalized eigenvalue problems presented in this paper allowed us to evaluate the proportions of the modes resulting from the discretization of the shallow water equations. The results obtained in section 4, show that the regular variation of the bathymetry does not influence the number of spurious inertial modes. We therefore retain that in the case of a constant bathymetry and in the case of bathymetries with regular variations, the proportions of the inertial spurious modes for the P_2-P_1 , $P_1^{NC}-P_1$, $P_1^{DG}-P_1$, $P_1^{DG}-P_2$ and P_0-P_1 finite elements pairs are 2/3, 4/7, 10/13, 1/4, 2/5, respectively. This study then makes it possible to generalize some results of Table 5 of (Le ROUX, 2012) to the case of bathymetries with regular variations.

The numerical results obtained from the discrete linear equations resulting from the formulations F1 and F2 show that the interaction between the approximation spaces of velocity field and the free surface elevation is the real cause of the appearance of spurious inertial modes. This is why in order to eliminate these spurious inertial modes we carry out a discretization in space of the 2D shallow water equations by using the discontinuous Galerkin method that approaches the velocity field and the free surface elevation in the same approximation space. We then determine the eigenvalues and the eigenvectors of the discrete linear system and we analyze these eigenvalues that correspond to the different modes of the system and we observe the absence of spurious inertial modes.

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