

Characteristics of Fluid Inclusion and D-O-S-Pb Isotopes Geochemistry of Jiaojia Gold Deposit

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Abstract

The Jiaojia gold deposit is located in the northwest jiaodong ore-concentrated area, north wing of Laizhou-Xixia anticlinoria, the orebody occurred in the granite which exists the footwall of Jiaojia Fault, Jiaojia Fault is the eastern secondary NE-NNE trending fault of Tan-Lu Great Fault. There are three types of fluid inclusion in Jiaojia gold deposit, respectively are H₂O-CO₂ inclusion (H-C type), CO₂-H₂O-CH₄ inclusion (CH₄-bearing type) and H₂O inclusion (H type). In Jiaojia gold deposit, the main ore-forming temperature is 170.2~358.1°C, the salinity is ranging from 2.04 wt% NaCl_{eq} to 10 wt% NaCl_{eq} with a average of 6.64 wt% NaCl_{eq}, sum intensity is (0.89~0.946) g/cm³, pressure estimation is 2×10³~3.8×10³kPa, the Jiaojia gold deposit is mesothermal gold deposit. The metallogenic substances derived from a mantle-crust-mixed source according to the S and Pb isotopes. The Jiaojia gold deposit was formed in the late stage which was a tensional environment. The ore-forming fluids mainly derived from magmatic and meteoric water, and a fraction of metamorphic water. The Jiaojia gold deposit is formed in the Yanshanian.

Keywords: Jiaojia gold deposit, fluid inclusion, isotope

1. Introduction

Jiaodong is one of the most important gold producing area in China, it reserves account for 1/3 of rock gold in our country has proven, its importance is self-evident. The Jiaojia gold deposit is located in the northwest Jiaodong ore-concentrated area, north wing of Laizhou-Xixia anticlinoria, the orebody occurred in the granite which exists the footwall of Jiaojia Fault, Jiaojia Fault is the eastern secondary NE-NNE trending fault of Tan-Lu Great Fault. Jiaojia gold deposit is one of the last three decades an important prospecting achievements for our broken with altered rock type gold, and achieved fruitful results of the prospecting, research results can be described as "voluminous".

Shandong Jiaojia gold deposit is the typical area for "Jiaojia Type" of broken belt altered rock type gold deposit, is also one of the important ore deposits in the geology study. However, the previous research of Jiaojia gold deposit is relatively inadequate and especially lack of systematic comprehensive geochemistry study such as fluid inclusion, stable isotopes, etc. therefore, it's important for us to study the geology characteristics, fluid inclusion and stable isotopes in this area, in order to conclude ore-forming regulations and metallogenic prediction in this region.

2. Deposit Geology

Jiaojia Gold deposit is located in the northwest of Jiaobei uplift, the north wing of Laizhou-Qixia anticlinorium, the secondary fault on the east of Yishu fault, occurred in the granite which formed in the footwall of Jiaojia main fault zone. The exposure strata in the regional context is dominated by the Archaean Jiaodong Group and also has the distribution of Proterozoic Fenzishan Group, Jingshan Group, upper proterozoic Penglai Group, mesozoic Group upper Jurassic and cretaceous system in local area, and the quaternary sediments is widely distributed (Fig. 1).

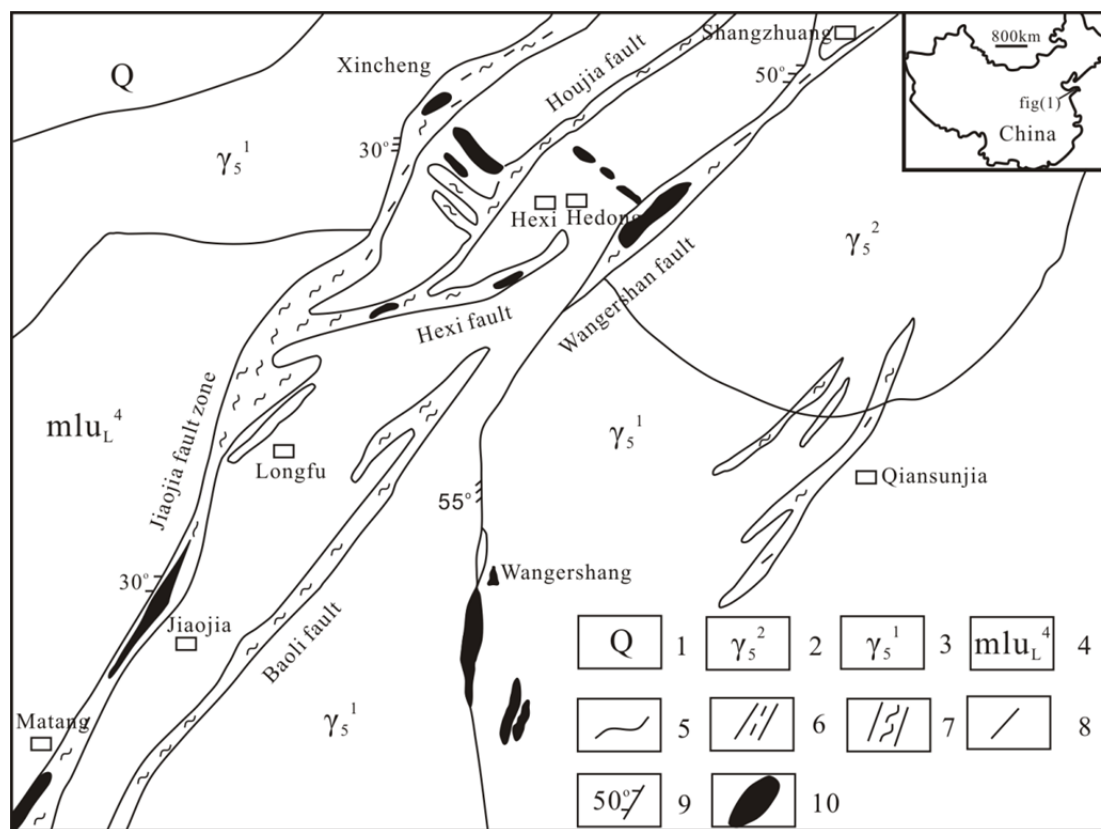


Figure 1. Geological map of Jiaojia gold deposit (modified after Li et al., 2002)

Note. 1. Quaternary; 2. Granodiorite of Guojialing; 3. Biotite granite of Linglong; 4. Metagabbro super unit of Malianzhuang (Amphibolite) ; 5. Geological boundary; 6. Early ore-controlling fracture alteration zone; 7. Middle mineralization alteration zone; 8. Fracture after mineralization; 9. The fracture and occurrence; 10. Gold orebodies.

The Jiaojia gold deposit is located in the middle section of Jiaojia Fault Belt where quaternary strata are extensively distributed. Bordered by main fracture, to the east is the Linglong monzonitic granite and Guojialing porphyritic granodiorite, to the west is the Malianzhuang metagabbro. The strata in the deposit area are simple, chiefly exposing Jiaodong Group of Archean Eonothem and thick Quaternary river alluvium, and major ore-controlling structure therein is Jiaojia fault. Demarcated by principal section, its eastern area is dominated by Linglong granite and western area is controlled by Jiaodong Group metamorphic strata. Jiaojia gold deposit occurs in the cataclastic alteration zone of Jiaojia principal section footwall. Besides, large-scale faults in this deposit area also include Hedong-Wang'shan fault, Sizhuang branch fault, Houjia fault, Baoli fault, etc. Exposed plutons here are Linglong gneissic biotite granite and Guojialing granodiorite. 80% of the whole area is dominated by magmatite, the dykes are well-developed as well (Fig. 1).

3. Stable Isotopes Geochemistry

3.1 Petrographic Characteristics of Fluid Inclusion

The type of fluid inclusion of Jiaojia Gold Mine mainly includes the H₂O-CO₂ inclusion (H-C type), CO₂-H₂O-CH₄ inclusion (including CH₄ type) and H₂O solution inclusion (H type) under the microscope. The H-C type is the main inclusion type, it presents 2 to 3 phases (H₂O liquid phase L_{H2O}+ CO₂ gas phase V_{CO2} and H₂O liquid phase L_{H2O}+CO₂ liquid phase L_{CO2}-CO₂ gas phase V_{CO2}) under the room temperature; and 2-phase is more common. The inclusion V_{CO2}+L_{CO2} generally takes up 30%-56% in volume of the inclusion; its size is mainly between 5-10μm, and the shape is oval, strip or irregular (Fig. 2-a, b). The H₂O-CO₂ inclusion is darker integrally, the boundary between the gas phase and liquid phase is blacker and thicker; the L_{CO2} (CO₂ liquid phase) is in matt dark color, and only the center has lighter points. At the early metallogenic stage, the H₂O-CO₂ inclusion is generated from tiny protogenetic inclusion; and it is the isolated or random distributed protogenetic inclusion at the main metallogenic stage. The inclusion (including CH₄) is mainly concentrated in the main

metallogenic epoch; the quantity is not so much; it coexists with the H-C type inclusion, and the composition is $\text{CO}_2\text{-H}_2\text{O}$ ($\pm\text{CH}_4$). The $\text{CO}_2\text{-H}_2\text{O}$ ($\pm\text{CH}_4$) inclusion presents 2 to 3 phases (H_2O liquid phase $\text{L}_{\text{H}_2\text{O}}$ + CO_2 liquid phase L_{CO_2} - CO_2 gas phase V_{CO_2} - CH_4) under the room temperature, wherein $\text{V}_{\text{CO}_2}+\text{L}_{\text{CO}_2}$ generally takes up 50%-70% of the total volume of inclusion (Fig. 2-c,d), with the size of 5-16 μm and in the oval or irregular shape. Most of inclusions are in matt dark color, and the boundary between gas phase and liquid phase is thick and black. This inclusion is rare at the early metallogenic stage, the main metallogenic epoch is represented as the random distributed protogenetic inclusion or distributed in the inclusion group, the random distributed protogenetic inclusion has larger gas liquid ratio change, and it often coexists with the H_2O solution inclusion. There is no obvious petrographic difference between the $\text{CO}_2\text{-H}_2\text{O}$ ($\pm\text{CH}_4$) inclusion and the $\text{H}_2\text{O-CO}_2$ inclusion under the microscope. According to the temperature measurement result, the complete melting temperature of solid CO_2 is lower than the pure CO_2 three-phase point (-56.6°C), but the existence of CH_4 may reduce $T_{\text{m,CO}_2}$ of inclusion, so it can be presumed that the $\text{CO}_2\text{-H}_2\text{O}$ ($\pm\text{CH}_4$) inclusion may appear in three phases. Such inclusion hardly exists at the late metallogenic epoch.

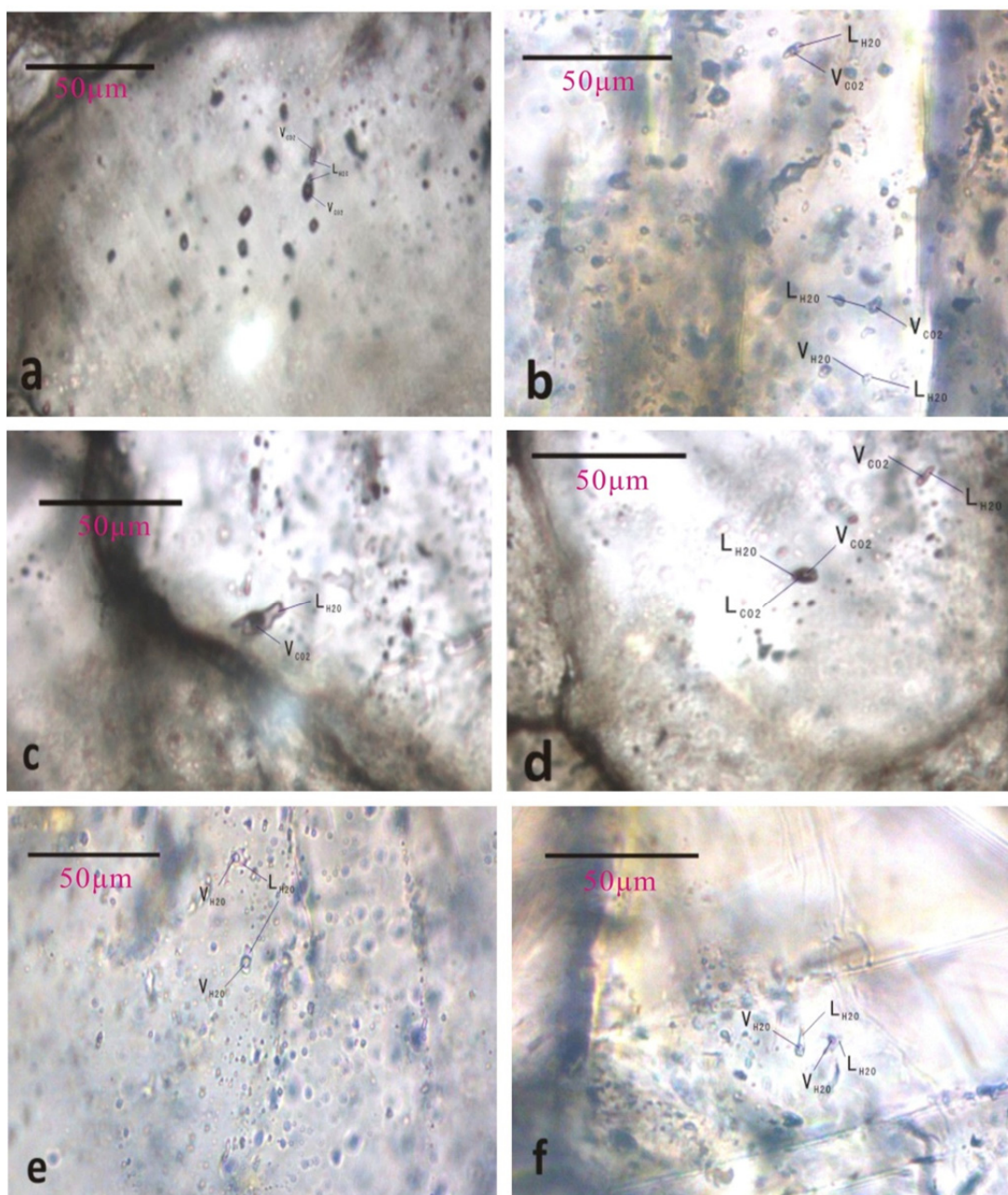


Figure 2. Fluid inclusion micrograph of Jiaojia gold deposit

Note. a. Lentoid H₂O-CO₂ inclusion is distributed in a scattered mode; b. Strip-type irregular H₂O-CO₂ inclusion and H₂O solution inclusion; c. Larger CO₂-H₂O-CH₄ inclusion; d. CO₂-H₂O-CH₄ three-phase inclusion and H₂O-CO₂ inclusion; (e, f) . irregular H₂O solution inclusion.

3.2 Microscopic Measurement Results

The fluid inclusion microthermometry analysis is conducted on the Linkan THMS 600 type heating-cooling stage of Chang'an University Metallogeny and Metallogenic Dynamics Key Fluid Inclusion Laboratory of the Ministry of Education, and temperature calibration is conducted on the heating-cooling stage by utilizing the artificially synthesized fluid inclusion standard sample provided by the U.S. FLUIDINC Company. The measurement precision of heating-cooling stage is $\pm 0.5^{\circ}\text{C}$ in the temperature interval of $-120^{\circ}\text{C} \sim -70^{\circ}\text{C}$, $\pm 0.2^{\circ}\text{C}$ in the interval of $-70 \sim +100^{\circ}\text{C}$ and $\pm 2^{\circ}\text{C}$ in the interval of $100 \sim 500^{\circ}\text{C}$. In the test process of fluid inclusion, the heating rate is generally $5 \sim 10^{\circ}\text{C}/\text{min}$, and it will be reduced to $2^{\circ}\text{C}/\text{min}$ when being close to the phase transformation point temperature so as to accurately record the transformation temperature of phase.

Through the detailed microthermometry on fluid inclusions of the representative sample at all metallogenic stages of Jiaojia Gold deposit and by combining with the inclusion study of predecessors on Jiaojia Gold Mine, it is concluded that the temperature range of main metallogenic epoch is between $170.2 \sim 358.1^{\circ}\text{C}$, the salinity (NaCl) is $2.04\% \sim 10\%$ with a mean value of $6.64\%(\text{NaCl})$.

3.3 Hydrogen and Oxygen Isotopes Geochemistry

The former researchers have conducted the hydrogen and oxygen isotope test on the inclusion samples in quartz and pyrite of deposit at different metallogenic stages, and the analysis result is at the Table 1. The analysis shows that the range of oxygen isotope values of quartz is between $10.4\% \sim 15.1\%$, which is lower than that of pyrite ($\delta^{18}\text{O}_{\text{V-SMOW}} = 12.04\% \sim 16.28\%$), and the hydrogen isotope of quartz is between $\delta\text{D}_{\text{V-SMOW}} = -32.99\% \sim -107\%$, which is higher than that of pyrite ($-83.68\% \sim -116.96\%$).

We projected the hydrogen and oxygen isotopes onto the diagram $\delta^{18}\text{O}-\delta\text{D}$ (Fig. 3), the projection points mainly fall in the range of magmatic water and atmospheric precipitation, and some points fall in the metamorphic water. The hydrogen and oxygen isotopes in the deep and shallow ore body are very similar. Therefore, the comprehensive analysis shows that the metallogenic hydrothermal of Jiaojia Gold Deposit is dominated by the atmospheric precipitation and magmatic water. Moreover, considering the existence of metamorphic water in the metallogenic process and the deep source of metallogenic material, it can be considered that the regional Jiaojia deep fracture plays an important role in the supply of hydrothermal solution during the mineralization.

Table 1. $\delta\text{D}-^{18}\text{O}$ isotope compositions in the Jiaojia gold deposit

Measured mineral	$\delta\text{D}_{\text{V-SMOW}}$	$\delta\text{O}_{\text{V-SMOW}}$	$\delta^{18}\text{D}_{\text{H}_2\text{O}}$	Data sources
Quartz	-76.6	14	6.14	Zhang et al., 2012
Quartz	-70.7	13.6	5.74	
Quartz	-69	15.1	7.24	
Quartz	-74.1	10.8	2.94	
Fluid inclusions in pyrite	-92.76	13.24	4.95	Song et al., 2013
Fluid inclusions in pyrite	-116.96	16.28	5.32	
Fluid inclusions in pyrite	-87.15	12.45	4.11	
Fluid inclusions in pyrite	-90.16	12.88	8.35	
Fluid inclusions in pyrite	-83.68	12.04	5.11	
Fluid inclusions in pyrite	-91.38	13.04	4.52	
Gold quartz polymetallic sulfide vein	-68.3	14.58	8.38	Xin, 2005
Calcite-quartz contains	-107	13.7	2.8	
While quartz contains	-75.4	10.4	3.94	Xu et al., 1997
While quartz contains	-32.99	13.74	9.69	

While quartz contains	-88.9	13.55	4.59	
While quartz contains	-95.8	13.5	4.54	
Gold - silver - quartz polymetallic sulfide vein	-75.6	13.97	5.01	
Ser	-62	10.4	7.25	Mao et al.,2008
Ser	-59	10.9	7.75	
Ser	-59	8.4	5.25	
K-feld	-86	10.2	4.37	
K-feld	-85	10	4.17	
Quartz	-107	12.55	2.4	Zhang et al.,1994
Auriferous pyrite	-89		-8.5	Lu et al., 1999
Quartz	-60.1		5.79	Zhang et al., 1989
Quartz	-60.1		6.69	Zhang et al., 1993

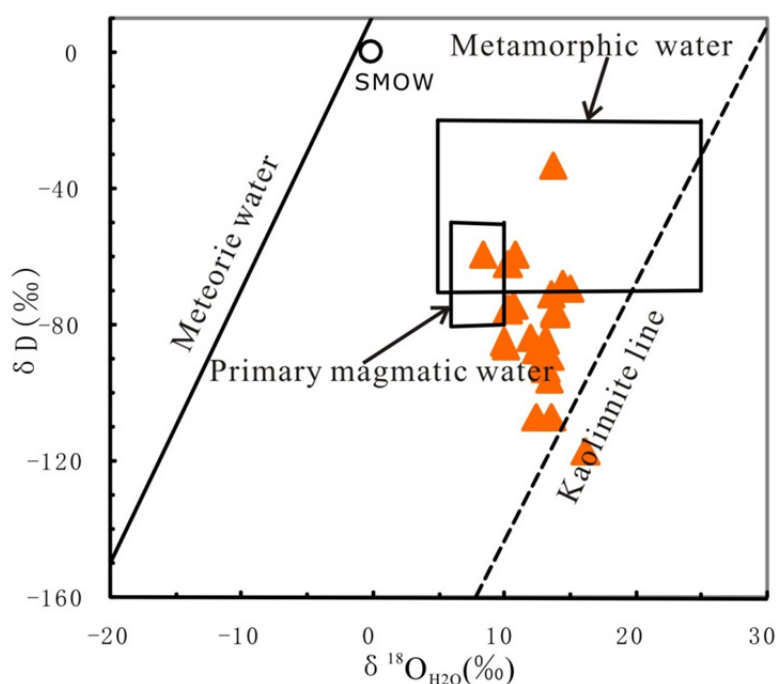


Figure 3. D–O isotopic composition diagrams of fluid inclusions of Jiaojia gold deposit

3.4 Sulfur Isotopes Geochemistry

The predecessors have obtained a lot of data of the sulfur isotope, which is shown in Table 2. According to the data in Table 2, we find that the sulfur isotope in the mining area is rich in $\delta^{34}\text{S}\text{‰}$, wherein $\delta^{34}\text{S}$ is between 6.5‰~12.72‰ with an average value of 10.46‰, these values are basically consistent with the $\delta^{34}\text{S}$ data (6.5‰~9.7‰, 8.03‰ in average; 10.1‰~12.2‰, 11.3‰ in average) proposed by Zhang (2012) and Mao et al. (2008), and it belongs to the positive deviation meteorite sulfur. Song et al. (2013) have tested the sulfur isotope of pyrite in Jiaojia deep gold deposit and found that the sulfur isotope of pyrite is relative stable, the variation range is small ($\delta^{34}\text{S}=11.08\text{‰}\sim 12.58\text{‰}$), and the sulfur isotope of the deep gold deposit is higher than that of the shallow gold deposit and located at the high value end in the sulfur isotope range (8.6‰~11.3‰) of shallow gold deposit.

Yang et al. (1998) conducted the sulfur isotope test on the main rock unit in the area, wherein the value of $\delta^{34}\text{S}$ value of Neoproterozoic amphibolites is 5.4‰ in average, and the $\delta^{34}\text{S}$ value of Yanshanian Linglong granite is 8.9‰ in average. The sulfur isotope of Jiaojia gold ore is distributed in the sulfur isotope range of surrounding

granite and Precambrian metamorphic series, but most are located at the high value portion, and it indicates that the source of sulfur in gold deposit is not the pure mantle sulfur or crust sulfur, but belongs to the mixed source. The magma of mantle source is contaminated with the crust substance formed by Precambrian metamorphic rocks and may be mixed with seawater, so formed a mixed sulfur which is rich in $\delta^{34}\text{S}$.

Table 2. Sulfur isotopic compositions of the Jiaojia gold deposits

Measured mineral	$\delta^{34}\text{S}(\text{‰})$	Data sources
Pyrite	11.22	Pang, 2005
Pyrite	10.99	
Pyrite	11.06	
Pyrite	10.75	
Pyrite	11.53	
Pyrite	10.92	
Pyrite	11.27	
Pyrite	10.79	
Pyrite	10.85	
Pyrite	10.59	
Pyrite	11.01	Zhang, 2012
Pyrite	8.3	
Pyrite	9.7	
Pyrite	8.4	
Pyrite	8.8	
Pyrite	7.9	
Pyrite	7	
Pyrite	6.5	
Pyrite	7.1	
Pyrite	8.6	Wang et al., 2001; Wang et al., 1991; Ding et al., 1998; Lin et al., 1999; Mao et al., 2008
Sulfide	10.3	
Pyrite	12.2	
Pyrite	11.4	
Pyrite	11.5	
Pyrite	11.3	
Pyrite	10.1	
Pyrite	11.12	Song et al., 2013
Pyrite	12.58	
Pyrite	11.46	
Pyrite	11.08	
Pyrite	11.23	
Pyrite	11.4	
Pyrite	11.36	
Pyrite	11.18	M et al., 2015
Pyrite	10.65	
Pyrite	11.74	
Pyrite	11.29	
Pyrite	8.48	
Pyrite	8.68	
Pyrite	10.63	
Pyrite	12.71	
Pyrite	8.82	
Pyrite	11	
Pyrite	10.26	
Pyrite	11.82	
Pyrite	12.72	
Pyrite	10.47	
Pyrite	11.16	

Pyrite	12.09	
Sulfide	9.65	Wang & Wang, 1991
Pyrite	10.27	
Pyrite	10.3	Wang, 2002
Pyrite	10.4	Wang & Wang, 1991
Pyrite	10.7	6th Geological Brigade of Shandong Province
Pyrite	10.2	Wang and Third Geological Brigade of Shandong Province
Pyrite	10.27	Yang et al., 1998

3.5 Lead Isotope Geochemistry

There are a lot of data about the characterization study of lead isotope obtained by former researchers (Tab. 3). According to Table 3, the lead isotope composition of Jiaojia Gold Deposit is relatively consistent ($^{206}\text{Pb}/^{204}\text{Pb}$ =17.117~17.37, $^{207}\text{Pb}/^{204}\text{Pb}$ =15.34~16.438 and $^{208}\text{Pb}/^{204}\text{Pb}$ =37.597~38.072). When projecting the data result on the $^{207}\text{Pb}/^{204}\text{Pb}$ - $^{206}\text{Pb}/^{204}\text{Pb}$ tectonic pattern diagram (Fig. 4-a), we

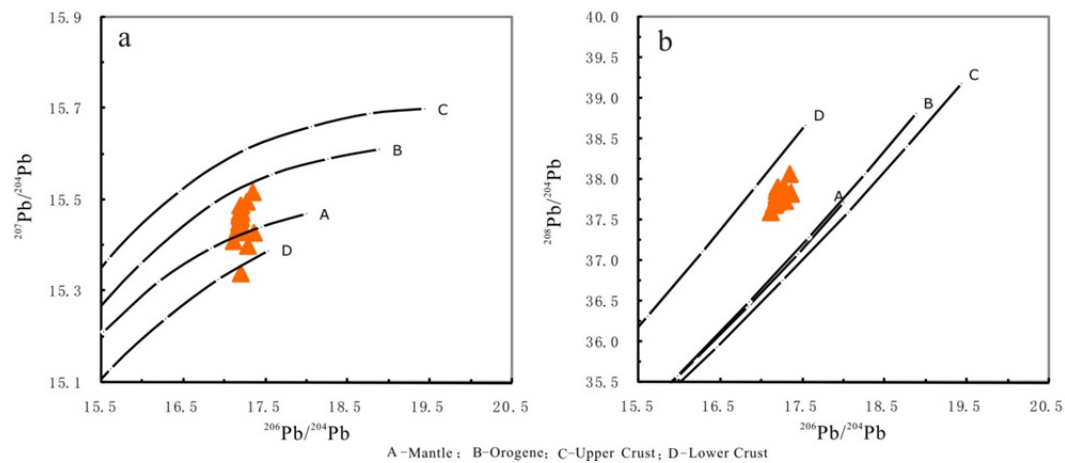


Figure 4. Discrimination diagram of Pb isotope composition for tectonic settings of Jiaojia gold deposit

can clearly see that the projection points mainly fall between the mantle and orogenic belt, and some points fall between the mantle and lower crust evolving curve; and in the $^{208}\text{Pb}/^{204}\text{Pb}$ - $^{206}\text{Pb}/^{204}\text{Pb}$ lead isotope tectonic pattern diagram (Fig. 4-b), the lead isotope projection points are very concentrated and fall between the orogenic belt and lower crust evolving curve. The above tectonic figures shows that the metallogenic material is dominated by the crust lead and mixed with mantle lead, and mineralization is formed in the tensional tectonic environment of relatively late period.

Regionally, the $^{206}\text{Pb}/^{204}\text{Pb}$ value of two ore-bearing rocks Linglong granite and amphibolites of Jiaojia gold deposit are respectively 17.320 and 17.233, the $^{207}\text{Pb}/^{204}\text{Pb}$ are respectively 15.415 and 15.474, and the $^{208}\text{Pb}/^{204}\text{Pb}$ are respectively 37.752 and 37.356 (Yang et al., 1996). The lead isotope ratio of two rocks is small in difference and shows regular changes, it shows that both rocks have the inheritance relationship on evolution, that is, the Linglong granite is formed by amphibolites under the action of magma remelting. The lead isotope ratio of ore is very close to the value of Precambrian metamorphic series, so it is predicted that the source of metallogenic material is mainly the mixture of ore-bearing granite and Precambrian metamorphic rocks.

Table 3. Lead isotopic compositions of Jiaojia gold deposits

Lead isotopic compositions of Jiaojia gold deposits			
$^{206}\text{Pb}/^{204}\text{Pb}$	$^{207}\text{Pb}/^{204}\text{Pb}$	$^{208}\text{Pb}/^{204}\text{Pb}$	Data sources
17.191	15.473	37.867	Yao et al., 1990
17.29	15.4	37.74	
17.203	15.462	37.815	

17.21	15.34	37.86	
17.353	15.518	38.072	6th Geological Brigade of Shandong Province
17.37	15.428	37.828	
17.167	15.43	37.683	Pang, 2005
17.2	15.451	37.752	
17.202	15.447	37.744	
17.212	15.489	37.908	
17.194	15.444	37.768	
17.212	15.472	37.825	
17.194	15.45	37.751	
17.203	15.477	37.855	
17.195	15.465	37.813	
17.191	15.467	37.796	
17.184	15.438	37.718	
17.203	16.162	37.815	Beijing Institute of Uranium Geology, Ministry of Nuclear Industry, 1983
17.21	15.34	37.89	Institute of Geochemistry, Chinese Academy of Sciences, 1981
17.37	16.438	37.828	Institute of Geochemistry, Chinese Academy of Sciences, 1987
17.25	15.43	37.82	Wang et al., 2001; Wang et al., 1991; Ding et al., 1998; Lin et al., 1999; Wen et al., 1990; Yao et al., 1990
17.203	15.46	37.815	Yao et al., 1983
15.21	15.34	37.89	Shenyang Institute of Geology and Mineral Resources
17.198	15.438	37.742	Ding et al., 1998
17.207	15.444	37.727	
17.117	15.41	37.597	
17.278	15.497	37.909	

4. Mineralogenetic Epoch

For the mineralogenetic epoch and granite study in Jiaodong area, the predecessors have accumulated a lot of data (As shown in Fig. 5), including the granite petrology and geochemistry (Li et al., 1986; Sang et al., 1984, 1992; Wang et al., 1985, 1986, 1988; Qiu et al., 1988; Wang et al., 1988; Xu et al., 1989; Yao et al., 1990; Li et al., 1993; Yang et al., 1996), granite factor (Qiu et al., 1988; Xu et al., 1989; Lv et al., 1993; Chen et al., 1993; Luo et al., 2002; Zhao et al., 2003), granite chronology (Luo et al., 2002) and granite genetic mineralogy (Chen et al., 1993). The Mesozoic granite distribution range in Jiaodong area is relatively broad, such as the “Linglong” biotite granite or gneissic granite in west of Jiaodong and the “Guajialing” maculosus granodiorite. Moreover, Kunyu Mountain- Que Mountain granite and granitic rock mass of Jiazi Mountain, Cha Mountain and Weide Mountain are developed in the east of Jiaodong (Guo, 2009). For the granite, the Linglong- Kunyu Mountain gneissic (biotite) monzonitic granite with the diagenesis age of 160-150Ma, the Guojialing maculosus granodiorite with the emplacement age of 130-126Ma and the Luanjia River-Weide Mountain coarse grain monzonitic granite with the age range of 127-105Ma have close relationship with Jiaodong gold mineralization (Lv, 2011). After long magma activity period of Mesozoic, the famous Jiaodong gold metallogenic epoch occurs, wherein the Mesozoic granite has a great relationship with Jiaodong gold mineralization, especially the Weide Mountain granite, it plays a crucial role in mineralization (Song et al., 2010). There are different opinions about the metallogenic epoch of Jiaodong Gold, for example, some scholars think that the metallogenic epoch is about 115±15Ma (Liu, 2001; Guo, 2009), 122.7~121.7Ma belongs to the Middle Yanshanian (Li, 1993) and 121.3-113.5Ma belongs to the Late Yanshanian (Song, 2011). Moreover, some scholars think that the Jiaodong area has two gold ores such as the gold ore of about 120-100Ma and the gold-silver-polymetallic ore of 80-90Ma (Yang, 2006), and the metallogenic age is mainly distributed in 120Ma. Due to the natural environment, the Sanshan Island gold ore and the Jiaojia gold ore field have the special and irreplaceable metallogenic geological environment (Li, 2013).

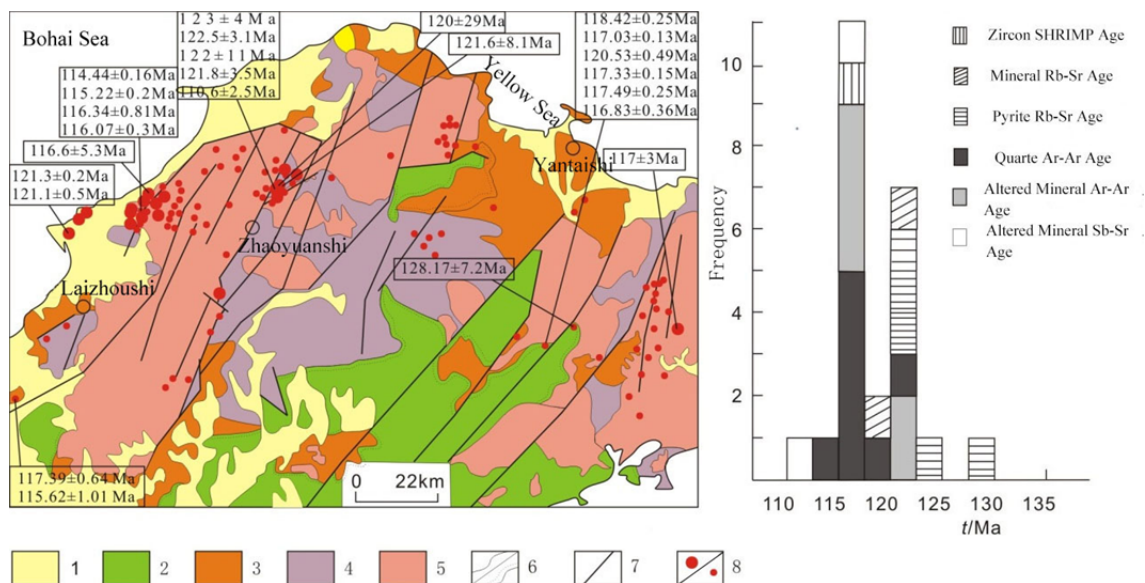


Figure 5. Spatial distribution and temporal histogram of gold deposition in the Jiaodong Peninsula (modified after Song et al., 2015)

Note. 1. Neogene+Paleogene; 2. Cretaceous; 3. Palenproterozoic and Neoproterozoic; 4. Precambrian granitic gneiss; 5. Jurassic-Cretaceous granite; 6. Conformable/unconformable geological boundary; 7. Fault; 8. Large-superlarge gold deposit/minor gold deposit.

5. Metallogenic Model

In the early and middle of Mesozoic, the paleo-pacific plate drove to the Eurasian plate rapidly from the SE direction to the NW inclined direction. Jiaodong area is at the collision orogeny stage and generates a series of conjugated brittle-ductile fracture systems (Fractures such as Sanshan Island, Jiaojia- Xincheng and Zhaoping) and the steep dip brittle fracture systems in the NNE direction in the northwest of Jiaodong peninsula. At the relaxation stage during collision orogeny, a lot of granite magma was generated through the syntactic or remelting function and intrusion emplacement occurs, and most granite rocks were formed at the footwall of brittle-ductile fracture. Formed in the Proterozoic era and parallel to the Qixia anticlinorium, the archaic fracture structure with the EW direction also generated the inheritance activities in this period. These fractures together with the brittle-ductile fault control the formation and distribution of the granite and gold deposit of this area.

In late Mesozoic, it is a crucial period of transition of dynamics system in Eastern Shandong since the Middle Jurassic, namely the conversion system of the collision to extension in late Mesozoic, which is characterized by readjustment of tectonic structure after subduction or the transition period of backward arc extensional dynamics environment. The Earth's surface is formed depression after extension and detachment, the upheaval of upper mantle prompts lithospheric delamination thinning, deep faults generated or activated, and undercut Tanlu fracture with strong sinistral strike-slip. This period is also the most intense period of the lithospheric thinning in eastern China, mainly intra-continental extension and lithospheric thinning, and there is mantle upwelling. With deep faults cutting into the upper mantle, strong upwelling of deep thermal and young asthenosphere material, accompanied by a large-scale and mixed crust-mantle material, with the mantle magma intrusion, geothermal gradient continues to grow, and further generates a lot of heat, resulting in lower crust or the crust-mantle re-melting with molten magma produced massive invasion activities and formation of metallogenic fluids, which leads to continuous consumption of lithospheric mantle and thinning. It results in a strong upwelling of deep mantle materials and energy; the upper mantle ore elements (Au, Te, Cu, etc.) were taken out (Au, Te, Cu, etc.). These elements move upward through the multi-stage evolution of mantle plume, leading to a large-scale, long-range migration. Ore-forming fluids continuously extract the minerals in the surrounding rock during the rising process. Due to the tectonic nature (toughness → brittle-ductile → brittle) and the transition of the physical and chemical conditions of ore containing fluid, immiscible fluids become immiscible and phase-separated, resulting in a fracture or toughness shear zone mineralization superimposed parts gathered, forming a large number of hydrothermal lode gold deposits.

6. Conclusions

The study of fluid inclusions determined that the Jiaojia gold Deposit is a medium-temperature deposits; metallogenic fluid is $\text{CO}_2\text{-H}_2\text{O-NaCl}(\pm \text{CH}_4)$ system; the main metallogenic temperature is $170.2\sim 358.1^\circ\text{C}$, salinity (NaCl) is $2.04\%\sim 10\%$; the average salinity (NaCl) is 6.64% ; an overall density is $0.89 \sim 0.946\text{g/cm}^3$, and the pressure estimate is $2 \times 10^3 \sim 3.8 \times 10^3\text{kPa}$.

The analysis and research on hydrogen and oxygen isotopes in Jiaojia Gold Deposit show that the hydrogen and oxygen isotopes from deep and shallow ore body are very similar. Ore-forming hydrothermal solution of Jiaojia Gold Deposit is derived from meteoric and magmatic water, a small part of metamorphic water also involved in the ore-forming process, indicating that the mineralization is a long process with inheritance. Combined with deep source of metallogenic material, Jiaojia fault has a major role in the gold mineralization structurally, which is regional, vertical and deep in spacial. In the process of mineralization, the regional tectonic - Jiaojia fracture has critical impact on the meteoric water.

The study of sulfur and lead isotopes of Jiaojia Gold Deposit shows the sulfur of Jiaojia Gold Deposit is not simply from mantle or crust, but is mixed sulfur. Mantle-derived magma mixed by crustal material formed by Precambrian metamorphic rocks, and may be mixed with seawater, forming mixed sulfur rich in $\delta^{34}\text{S}$. The main source of lead derived from crust, and mixed with the mantle lead, mineralization formed in a relatively late extensional background. The ratio of lead isotope of ore and granite in surrounding rock is very close to Precambrian metamorphic rocks series, so we presume that the source rocks of mineralization are mainly ore-bearing granitoids and Precambrian metamorphic rocks.

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